

The Effect of Wind on Shoaling Wave Shape

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Onshore (Feddersen *et al.* in prep.)

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Wind and Ocean Waves

- Wind causes growth (Jeffreys 1925; Miles 1957; Phillips 1957)
- Pressure and wave phase difference pumps energy
- Simplest model for pressure p (Jeffreys 1925)

$$p \propto \frac{\partial \eta}{\partial x} \quad \text{at} \quad z = \eta$$

with η wave surface

- Growth know, shape change new (Zdyrski and Feddersen 2020, in JFM)

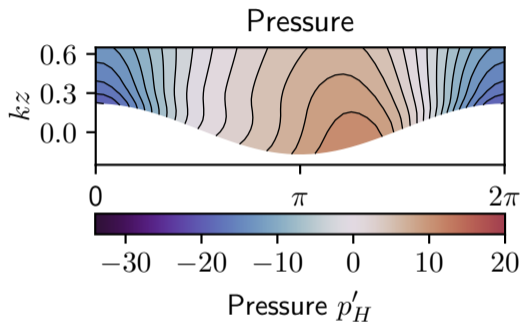


Figure 1: LES simulation of pressure above a wave (Husain et al. 2019).

Wind and Wave Shape

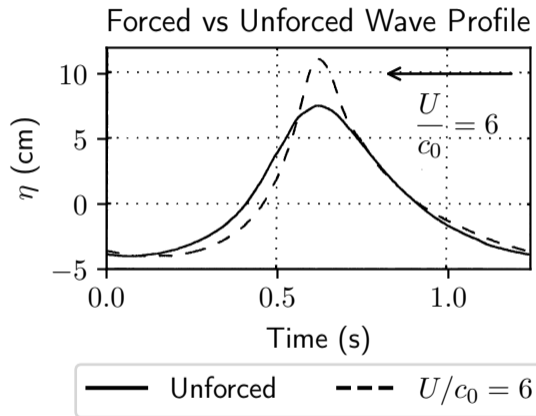


Figure 2: Reproduced from Feddersen and Veron (2005).

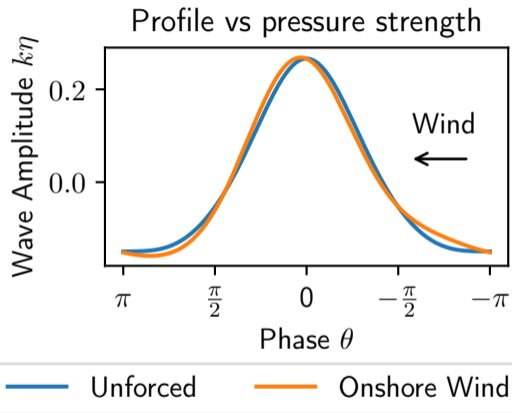


Figure 3: Reproduced from Zdyrski and Feddersen (2020).

Wave Shoaling

- Decreasing water depth (shoaling) also changes wave shape (Elgar and Guza 1985)
- Changes skewness S and asymmetry A

$$S = \frac{\langle \eta^3 \rangle}{\langle \eta^2 \rangle^{3/2}} \quad \text{and} \quad A = \frac{\langle \mathcal{H}[\eta]^3 \rangle}{\langle \mathcal{H}[\eta]^2 \rangle^{3/2}}$$

- $\langle \cdot \rangle$ average over wave period and \mathcal{H} Hilbert transform
- Goal: examine shape changes from wind and shoaling, applicable in the nearshore

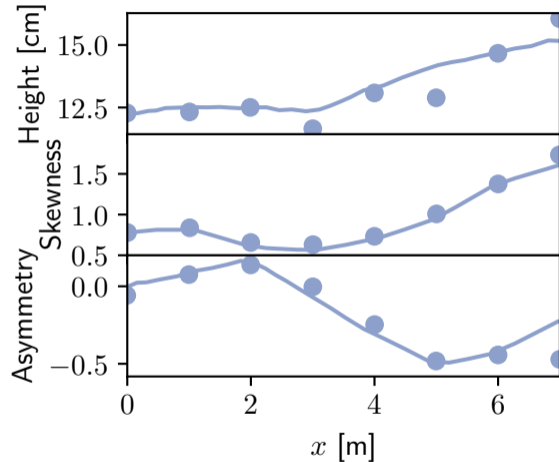


Figure 4: Shape statistics for shoaling, unforced wave experiment; reproduced from Cienfuegos, Barthélemy, and Bonneton (2010).

Setup

- Incompressible, irrotational, inviscid, 2D flow
- $\eta(x, t)$ and $\nabla\phi(x, t, z) = \vec{u}$
- Pressure enters Bernoulli equation

$$0 = g\eta + \frac{\partial\phi}{\partial t} + \frac{1}{2} \left[\left(\frac{\partial\phi}{\partial x} \right)^2 + \left(\frac{\partial\phi}{\partial z} \right)^2 \right] + \frac{p}{\rho_w} \quad \text{at } z = \eta$$

- Jeffreys forcing $p_J = P\partial_x\eta(x, t)$
- Assume $\partial_x h \ll 1$, so bottom BC is just $\partial_z\phi = 0$ to leading order
- Four free, nondimensional parameters:
 - a/h_0 (amplitude)
 - $\partial_x h$ (slope)
 - kh_0 (depth)
 - $Pk/(\rho_w g)$ (pressure magnitude)

- Assume $\varepsilon := a/h_0 \sim (kh_0)^2 \sim Pk/(\rho_w g) \ll 1$ and $\partial_x h \sim \varepsilon^{3/2}$
- Method of Multiple Scales
 - $\eta = \varepsilon\eta_1 + \varepsilon^2\eta_2 + \dots$
 - $x_0 = x, x_1 = \varepsilon x, x_2 = \varepsilon^2 x, \dots$
 - Co-moving coordinate $\xi_+ = -t + \int^{x_0} dx/c(x)$
- Variable-coefficient Korteweg-de Vries Burgers equation

$$\frac{\partial \eta_0}{\partial x_1} + \frac{3}{2} \frac{c_0^3}{c^3(x_1)} \frac{\eta_0}{a_0} \frac{\partial \eta_0}{\partial \xi_+} + \frac{1}{6} \frac{h^3}{a_0} \frac{c(x_1)}{c_0} \frac{\partial^3 \eta_0}{\partial \xi_+^3} + \frac{1}{2c} \frac{\partial c(x_1)}{\partial x_1} \eta_0 = -\frac{1}{2} \frac{P}{\rho_w g} \frac{c_0^2}{c^2(x_1)} \frac{\partial^2 \eta_0}{\partial \xi_+^2}.$$

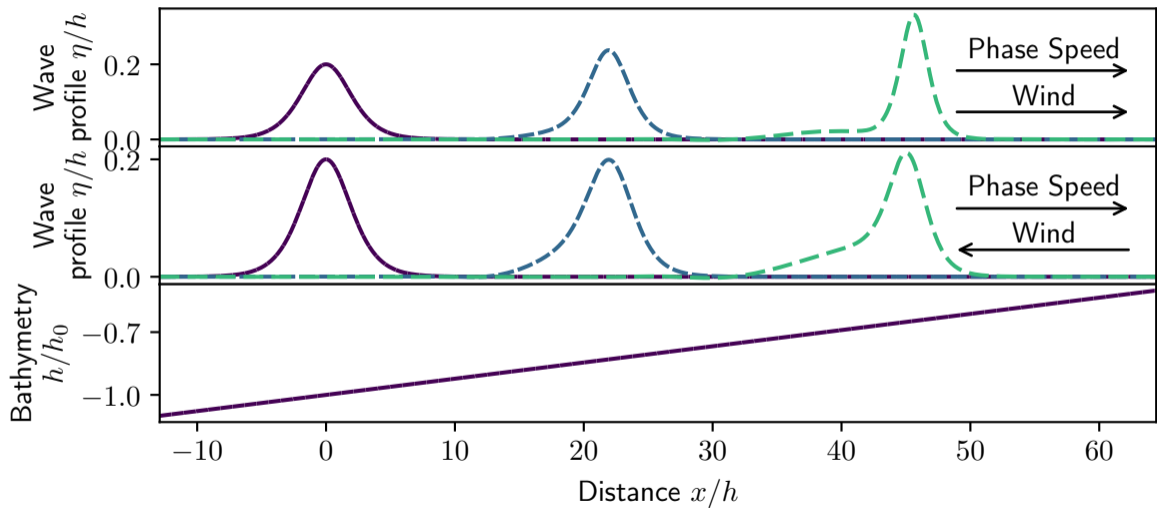
with $c(x_1) = \sqrt{gh(x_1)}$ and $c_0 = \sqrt{gh_0}$

- Solitary waves initial conditions

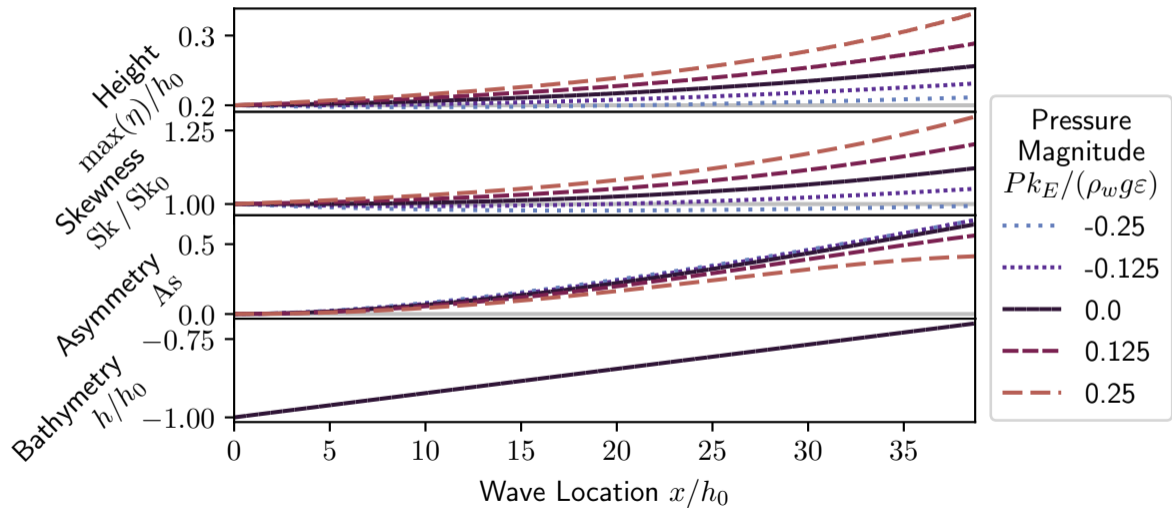
$$\eta_0 \Big|_{x_1=0} = 2 \operatorname{sech}^2 \left[\sqrt{\frac{3}{2}} \xi_+ \right]$$

- Sign of P depends on wind direction: onshore wind $\implies P > 0$ and growth
- Solve numerically with RK3 central difference scheme and $h(x_1)/h_0 = 1 - 0.1kx_1$

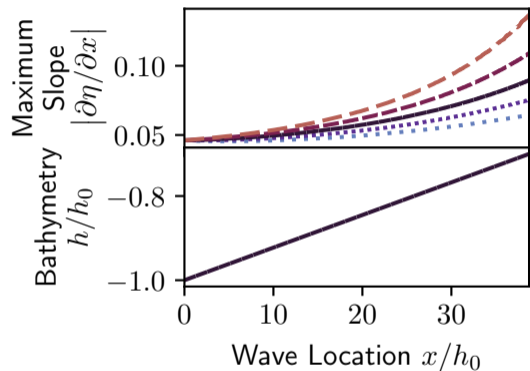
Results: Profile



Results: Height, Skewness, and Asymmetry



Results: Steepness and Breaking



Pressure Magnitude $Pk_E/(\rho_w g \epsilon)$

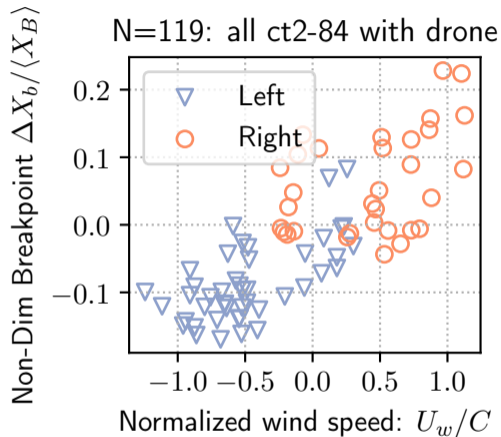
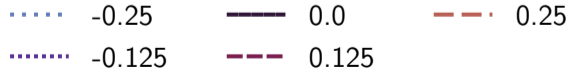


Figure 5: Break point data vs. wind speed collected at the Surf Ranch, CA (Feddersen *et al.* in prep.)

Discussion and Conclusion

- Coupled surface pressure to the Bernoulli equation with gently sloping bottom
- Method of Multiple Scales produced damped, variable-coefficient KdV-Burgers equation
- Numerically calculated shape changes consistent unforced shoaling waves
- Derived wind-induced height, skewness, and asymmetry
- **Onshore wind causes steepening in deeper water, while offshore wind delays steepening to shallower water**

Future Work:

- Extend results to periodic waves
- Include dynamic wind-wave coupling



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Part I

Appendix

Laplace Equation and Boundary Conditions

$$\frac{\partial^2 \phi}{\partial x^2} + \frac{\partial^2 \phi}{\partial z^2} = 0 \quad (1)$$

$$\left. \frac{\partial \phi}{\partial z} \right|_{z=-\infty} = 0 \quad (2)$$

$$\left. \frac{\partial \phi}{\partial z} \right|_{z=\eta} = \frac{\partial \eta}{\partial t} + \frac{\partial \eta}{\partial x} \left. \frac{\partial \phi}{\partial x} \right|_{z=\eta} \quad (3)$$

$$0 = g\eta + \left. \frac{\partial \phi}{\partial t} \right|_{z=\eta} + \frac{1}{2} \left(\left(\left. \frac{\partial \phi}{\partial x} \right|_{z=\eta} \right)^2 + \left(\left. \frac{\partial \phi}{\partial z} \right|_{z=\eta} \right)^2 \right) + \frac{p}{\rho_w g} \quad (4)$$

with

$$\vec{u} = \nabla \phi \quad (5)$$

Constraints:

- Periodic¹

$$\vec{u}(x, z, t) = \vec{u}(x + L, z, t)$$

- Progressive

$$\vec{u}(x, z, t) = \vec{u}'(x - \tau(t), z)$$

- No current $\langle \vec{u} \rangle = 0$

¹Note: this precludes sloping bottom topographies

Wave Slope

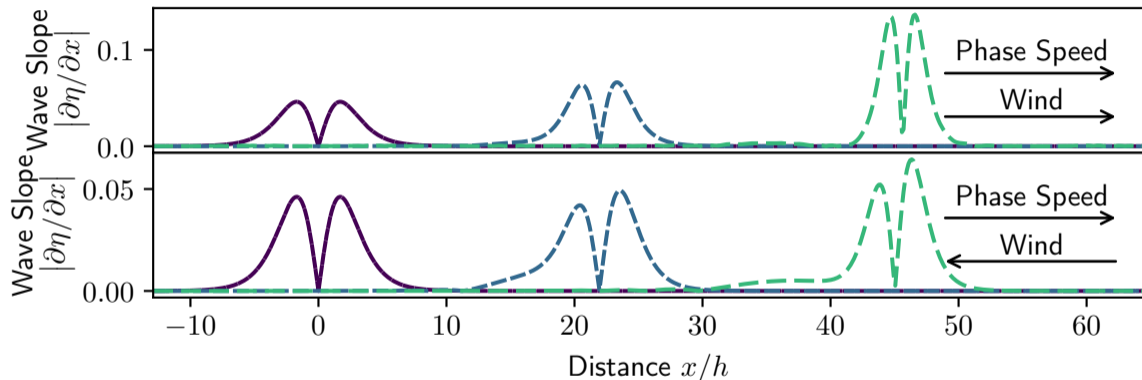


Figure 6: Magnitude of wave slope as a function of distance x/h_0 for (a) onshore and (b) offshore winds.